EARLY CRETACEOUS PALYNOMORPHS FROM THE ZUATA AREA, EASTERN VENEZUELA

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RESUMO²

Este trabilho é baseado em estudos palinológicos e petrográficos do testemunho de um popo (MARAVEN, S. A.), situados airs de a Cuatat, Nenzeucal. Os palinomorfos recuperados são atribuidos a 19 gêneros e 33 espécies de esporos e 9 gêneros e 26 espécies de polen. A idade indicada é, mas provavelment. Aptiano-Albano medito inferior. São feiras comparações com modelos ji estabelecados de provincias micenteristicas do Creticoo Interior (MRINNER, 1976, IPROVIGERA) bente de deposição fluvial jabo condições strifas, numa região toopparficamente pouco novimentada,

ABSTRACT

This report is based on palynological and petrographical studies of a core from a Zusta area well MARAYER, S. A.Palynomorphis encountered have been assigned to 19 genera and 38 pecies of spores and 40 genera and 26 species of polien grains. The most probable age is "Aptitus-Early Middle Malliss. A. comparison with previously published Early Cretaceous mercondroal province concepts of the property of the propert

INTRODUCTION

This report is based on palynological and pertographical studies of a core sample from a Zuata area well (MARAVEN, S.A.), drilled in the petroleum belt of the Orinoco River, Faja Petrolifera del Orinoco (Fig. 1). The core (Fig. 2) represents an interval of 20 (6.096m) from 4175' (1272.5m) to 4195' (1278.6m) within the basal section of the Cretacous Temblador Formation, which closely overlies the top of the Paleozoic basement (16' below).

The Early Cretaceous palynomorphs recovered from the sample at 4176'9" (1273.11m) belong to 19 genera and 33 species of spores and 9 genera and 26 species of pollen. Marine palynomorphs are absent.

This study, to the author's knowledge, is the first description of an Early Cretaceous microflora from Venezuela. Other studies of Cretaceous palynology in South America may be found in BOER et al. (1965), MULLER (1966), BRENNER (1968). REGALI. UESUGUI & SANTOS (1974), LIMA (1974) and HERN-GREEN (1972, 1973, 1974, 1975, 1976, 1981).

Here an attempt has been made to place the recorded assemblage into Benner's (1976) Barremian-Cenomanian floral provinces as well Hemgreen's (1976, 1981) Cretaceous microfloral provinces and to discuss the implications. It is hoped that the results obtained will provide valuable assistance to other researchers of South America and Africa and also give some insight into the early migration of anziosperms.

PETROGRAPHICAL DESCRIPTION OF THE CORE

Petrographical and Scanning Electron Microscopy (SEM) studies have shown that the lithology consists mainly of white, well-consolidated, detrital, kaolinite-rich (20-80%), moderately sorted medium- to coarse-grained santstone with subangular to subrounded, fractured

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Fig. 1 - Location map

quartz grains. The sedimentary structures consist of crudely stratified beds, interbedded with small scale cross-bedding (20-80 mm) which are often separated by subhorizon-tal erosional surfaces stained by reddish iron toxide minerals. Palynomorphs have been recovered from a knolinite-rich, light-grey, wellsorted, fine to very fine sandy (subangular to subrounded quartz grains) kaoliniti-cilitic claystone, rich in earbonaceous matter (410°–44777).

SAMPLE PREPARATION AND METHOD OF STUDY

Standard palynological and maceration procedure was used in preparation of the sample; namely, treatment with HCI, HF, centri-luging and the use of zinc bromined (specific gravity 2.2) as a heavy liquid; followed by oxidation by HRO, for a short period with brief clearance in alkali using ammonium ity-action of the control of the cont

The slides were examined at 400X magnification using a Zeiss photomicroscope equipped with a standard mechanical stage. Photographs were taken using Agfa Isopan 21 film. Representatives of all spore and pollen taxa encountered are illustrated in Pls. I - XVII. All slides are on file in the Biostratigraphy section of the Earth Sciences Department, INTEVEP, S.A.

COMPOSITION OF THE MICROFLORA

The palynologically productive sample at 4176'9" (1273.11m) has yielded 59 species of uncorroded and well-preserved spores and pollen

The assemblage, in general, is dominated by germooperm pollen and pierdophyte spores. Angiosperm pollen is extremely rare and its coccurrence does not exceed. 2% to 4% of the total assemblage. Bisaccates, tricolporates, triporties and polyporates have not been encountered. Broadly, it can be concluded that the microflors. Behedra-type pollen and several polyporates and the proposer of the proposers of the propo

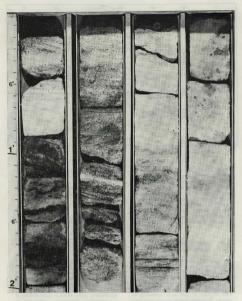


Fig. 2 — Part of the core (4,175' to 4,195' interval) through the basal section of the Temblador Formation, Quata area, Faja Petrolifera del Orinoco, Venezuela. Palynomorphs were recovered from the dark, carbonaceous portion in the lower left corner of the figure (4,176' to 4,177' interval).

assemblage. This can be illustrated by the occurrence of Cicatricosisporites spp. in relatively high values (15%). Other quantitatively important constituents among the pteridophytes are species of Cyathidites (8% to 10%) and Concavisporites (5% to 7%). Apart from the above--mentioned genera all others encountered do not seem to have any quantitative significance. The following species occur in percentages lower than 5%, some of these species being recorded only as single specimens. Gleicheniidites anilohatus Brenner 1963 (Plate II. Fig. 2). Verrucosisporites sp. A (Plate II, Fig. 3), Verrucosisporites sp. B (Plate II, Fig. 4), Biretisporites cf. B. notoniaei Delcourt and Sprumont 1955 (Plate II, Fig. 6), Dictyophyllidites equiexinus (Couper) Dettmann 1963 (Plate II, Fig. 7), Contignisporites fornicatus Dettmann 1963 (Plate VII, Fig. 3), Leptolepitides verrucatus Couper 1958 (Plate VIII, Fig. 1), Leptolepitides major Couper 1953 (Plate IX, Fig. 2), Foveotriletes sp. A (Plate VIII, Fig. 2), Klukisporites variegatus Couper 1958 (Plate VIII, Fig. 3), Lycopodiumsporites paleocenicus Dutta and Sah 1970 (Plate VIII, Fig. 4), Cingulatisporites verrucatus Regali, Uesugui and Santos 1974 (Plate IX. Fig. 3), Sergipea naviformis Regali, Uesugui and Santos 1974 (Plate IX. Fig. 5), Coptospora sp. A (Plate IX, Fig. 6), Coptospora sp. B (Plate IX, Fig. 7), Schizea reticulata Cookson 1956 (Plate XI, Fig. 1). Schizea certa Bolkhovitina 1956 (Plate XI. Fig. 2), Hamulatisporis hamulatis Krutzsch 1959 (Plate IX, Fig. 1), Velosporites sp. A (Plate X, Fig. 1), Callialasporites dampieri (Balme) Deu 1961 (Plate X, Fig. 2 and 4), Perotriletes pannuceus Brenner 1973 (Plate X. Fig. 3), Gnetaceaepollenites sp. A, (Plate XII, Fig. 1), Gnetaceaepollenites sp. B. (Plate XII. Fig. 2 and 4), Gnetaceaepollenites sp. C, (Plate XII, Fig. 3), Gnetaceaepollenites sp. D (Plate XII, Fig. 5), Gnetaceaepollenites sp. E (Plate XII, Fig. 6), Zonalapollenites sp. A (Plate XV, Fig. 4). Retitricolnites vulgaris Pierce 1961 (Plate XV. Fig. 5), Eucommidites sp. A (Plate XV. Fig. 6). Retimonocolpites sp. A (Plate XVIII, Fig. 4) and Monosulcites sp. A (Plate XVIII, Fig. 5).

DISCUSSION AND CONCLUSIONS

The concept of Barremian-Cenomanian floral provinces was introduced by BRENNER in 1976. The provinces he proposed from south to north (Fig. 3) and their possible climatic patterns are as follows:

The Northern Laurasian Province (NLP):
Temperate humid climate; now above

The Southern Laurasian Province (SLP): Warm temperate to subtropical humid; now in the middle latitudes of the Northern Hemisphere,

The Northern Gondwana Province (NGP):
Tropical semiarid; presently from the
northern coast of South America and Africa
to some undetermined latitude south of the
equator,

The Southern Gondwana Province (SGP):
Warm temperate to subtropical humid;
now in southern regions of South America and
Africa, and including Australia, New Zealand
and India.

According to Brenner's concept. Venezuela should be placed in NGP, and indeed the microflora does show some resemblance to Aptian-Albian assemblages from other localities in similar latitudes also placed within NGP. However some dissimilarities exist, Most notably this assemblage contains no bisaccate pollen, which is normally considered one of the most important characteristics of this province, Equally, while the low diversity of pteridophyte spores is characteristic, the Venezuelan material is somewhat unusual in the higher frequency of Schizeaceous types, whereas these have previously been thought to be rare (BRENNER, 1976). The absence of numerous and diverse cycadophytes is also anomalous, In northeastern Peru, Brazil, West Africa and Israel the polyplicate grains of the Ephedraceae are very numerous and diverse, commonly up to 10 to 20% of the assemblage, which is also considered another important characteristic of this province. This feature seems consistent with the Venezuelan material. Comparisons of Middle Cretaceous spores and pollen from West Africa as described by JARDINE & MAGLOIRE (1965) with the same age assemblages from northeastern Peru have been summarized by BRENNER (1968), During the Late Albian and Early Cenomanian both areas contained identical species of Galeacornea and Elaterocolpites and also perinate trilete types, These features were used to correlate the Upper Albian to Lower Cenomanian in both areas. In Venezuela, the presently recorded assemblage does not include these genera, and they are absent with the exception of a single specimen of Perotriletes pannuceus. The foregoing suggests that the material studied should be considered older than Late Albian.



distribution of palynoprovince during Barremia Cenomenian tim (taken from S NANOGLU FREMEN 1980

The concept of an African – South Amrrican Middle Createous microfloral Province (A.S.A), was introduced by HERNGREEN (1975) and extended in a contribution to the IV International Palymology Conferee, Lucknow, India (HERNGREEN, 1976). Later in 1981, after reviewing the great majority of course prince of the Province of the Control (1981) established the following microfloral provinces:

The Boreal Early Cretaceous province of the Northern Hemisphere.
 The Late Cretaceous Normapolles Pro-

vince.

3) The Late Cretaceous Aquilapollenites

Province.

4) The pre-Albian West African-South American Province (WASA).

5) The Middle Cretaceous (Albian-Cenomanian) African-South American Province

(ASA).

6) The Late Cretaceous Palmae Province of Africa and Northern South America.

7) The Gondwana Province.

8) The Senonian Nothofagidites Province.

The ASA phytogeographical belt is characterized by the presence of species bearing elaters and a rich morphological diversity of the ephedroid complex, high percentages and conspicuous diversification of the angiosperm pollen, the absence of bi- and trisaccate gymnosperms and the absence or scarcity of fern taxa. The Cenomanian (and Turonian to probably Conjacian) microfloras in the Sergipe (Brazil) and Congo Basins are dominated by Hexaporotricolpites. Based on these characteristics. HERNGREEN (1975) draws the southern boundary of the ASA floral province from Peru (and northern Chile?) to the Alagoas/ Sergipe basins into West Africa somewhere between the Congo and Angola to East Africa (Somalia). The northern boundary is drawn through the northernmost parts of South America and via Algeria and Tunisia to Israel. It is important and relevant to the present study that he considers the southern limit of the ASA microfloral province of a provisional nature.

Similarities include the presence of a great number and variety of the ephedroid complex and absence of bi- and trisaccate pollen. Dissimilarities include the absence of elater-bearing species. The lack of the high percentages and diversification of the angiosperm pollen in the recorded assemblage and the absence of Hexaporotricolpites genera are noteworthy. Taking these characteristics into account, the age of the presently recorded assemblage does not seem to be older than Albian. The palynomorph content of several borehole sections in the Barreirinhas and Sergipe basins of Brazil has been described and palynological zones have been established by HERNGREEN (1973, 1974, 1975). Later in 1975 he compiled and summarized his results with some additional information from several other boreholes. The palynological zones of the Barreirinhas basin are as follow:

I. The Reyrea polymorphus zone; Early-Middle Albian

IA. The Elateropollenites jardinei and Ephedripites irregularis subzone

IB. The Elaterosporites protensus and Elaterosporites verrucatus subzone II. The Elaterocolpites-Elateroplicites-So-

II. The Etaterocolpites-Etateropticites-Sofrepites zone; Late Albian-Early-Cenomanian IIA. The Classopollis spinosus subzone IIB. The Classopollis braziliensis subzone

III. The Triorites africaensis zone; Late Cenomanian

IIIA. This zone is not formalized
IIIB. The Ephedripites elsikii subzone

Striking similarities exist between the recorded assemblage and Herngreen's R. polymorphus zone, more specifically with the I A

subzone, except for the abence of R. polymorphus and E. familed. In fact HENGKEEN (1975) also mentions the infrequent occurrence of these species. Elater-bearing species were also not recorded from the Bahia or Sergipel. Alagoas basins (MULLER, 1966). The recorded assemblage does not show any similarities with the other younger zones of HENGKEEN (1975). Based on the above-mentioned cristic properties of the subsemblage is Early Middle Albhan.

In 'terms of paleoenvironmental conditions, petrographical studies suggest that these beds are the product of a fluvial environment, probably a point bar, which may account for the kasolinite-cich and medium-to coarse-grained the fluodipalion or the swamp during a short period and is recorded in the dark- to light gery, bioturbated and fine sandy kolinite-illite mixed claystone rich in plant remains. The source of the quartz sand and the detrital kasolinite-dilite could be from the adjournal Skield.

In terms of palynological interpretations, the absence of matrie palynomorphs, as well as the presence of Classopolis species in relatively high values, suggests a shallow water "sub-aerial" noumarine environment. The high percentagem of numerous Ephedripites species and the diversity of the pteridophyte spores further suggest and conditions. The total absence of bisaccate pollen indicates rather flat tonogramble.

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PLATES I-XVII

Editors' note: The original plates have been reproduced here four to a page as an economy measure. All figures are at the same magnification, except in Pt. XVII; according to the author, the bar scale in the figure in the upper left corner of each page represents Zbum and serves the entire page.

PLATE I

Figs. 1, 2 — Cyathidites minor Couper 1953. Fig. 3 — Cyathidites australis Couper 1953. Fig. 4 — Concavisporites sp. A.

PLATE II

Fig. 1 — Cyathidites sp. A. Fig. 2 — Gleicheniidites apilobstus Brenner 1963. Fig. 3 — Verrucosisporites sp. A. Fig. 4 — Verrucosisporites sp. B. Fig. 5 — Concevisporites sp. B. Fig. 6 — Biretisporites cf. 8. potoniaei Delcourt 8 Sormant 1965. Fig. 7 — Dictrophrilidites equisivius (Couper) Dettmann 1964.

PLATE III

Fig. 1 — Cicatricosisporites patamacensis Brenner 1963. Fig. 2 — Cicatricosisporites avnimelechi Horowitz 1970. Fig. 3 — Cicatricosisporites dorogenesis Potonié 1934.

PLATE IV

Fig. 1 — Cicetricosisporites patomacensis Brenner 1963. Fig. 2 — Cicetricosisporites avnimelechi Horowitz

PLATE V

Figs. 1, 2 - Cicatricosisporites australiansis (Cookson) Potonia 1965.

PLATE VI

Figs, 1, 2-Cicetricosisporites recticicatricosus Dorning 1965, Fig. 3-Cicetricosisporites subrotundus Brenner 1963, Fig. 4-Cicetricosisporites sp. A.

PLATE VII

Fig. 1 — Cicatricosisporites sp. B. Fig. 2 — Cicatricosisporites sp. C. Fig. 3 — Contignisporites fornicatus Dettmann 1963.

PLATE VIII

Fig. 1 — Leptolepitides verrucatus Couper 1958. Fig. 2 — Foveotriletes sp. A. Fig. 3 — Klukisporites variegatus Couper 1958. Fig. 4 — Lycopodiumsporites peleoocenicus Dutta & Sah 1970.

PLATE IX

Fig. 1 — Hamulatisporis hamulatis Krutzsch 1959, Fig. 2 — Leptolepidites major Couper 1953, Fig. 3 — Cingulatisporites verrucatus Regali, Uesugui & Santos 1974, Fig. 4 — Cingulatisporites sp. A. Fig. 5 — Sergipes naviformis Regali, Uesugui & Santos 1974, Fig. 6 — Coptopora sp. 8. — Fig. 7 — Coptopora sp. 8.

PLATE X

Fig. 1 — Velosporites sp. A. — Fig. 2 — Callielasporites dampieri (Balme) Dev 1961. Fig. 3 — Perotriletes pannuceus Brenner 1973.

PI ATE XI

Fig. 1 - Schizes reticulate Cookson 1956, Figs. 2, 3 - Schizes certs Bolkhovitina 1956.

PLATE XII

Fig. 1 – Gnetaceaepollenites sp. A. Figs. 2, 4 – Gnetaceaepollenites sp. B. Fig. 3 – Gnetaceaepollenites sp. C. Fig. 5 – Gnetaceaepollenites sp. D. Fig. 6 – Gnetaceaepollenites sp. E.

PLATE XIII

Fig. 1, 2 — Ephedripites strigatus Brenner 1968, Fig. 3 — Ephedripites sp. B. Fig. 4 — Ephedripites ambonoides Brenner 1968.

PLATE XIV

Fig. 1 — Ephedripites procerus Brenner 1968. Fig. 2 — Ephedripites berghoornii Pocok 1964. Figs. 3, 4 — Ephedripites sp. C.

PLATE XV

Fig. 1 – Ephedripites validus Brenner 1968. Fig. 2 – Ephedripites sp. D. Fig. 3 – Ephedripites sp. E. Fig. 4 – Zonelapollenites sp. A. Fig. 5 – Retitricolpites vulgaris Pierce 1961. Fig. 6 – Eucommidites sp. A.

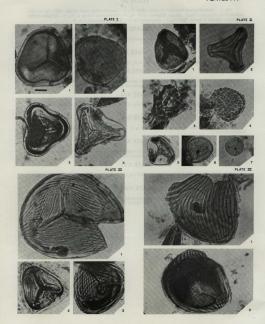
PLATE XVI

Figs. 1,3,5 — Classopollis jardinai Herngreen 1973. Fig. 2 — Classopollis sp. A. Figs. 4,8 — Classopollis obidosensis Groot & Groot 1962.

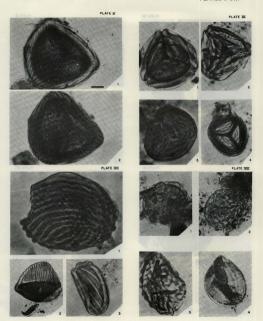
PLATE XVII

Figs. 1,3 — Classopoliis sp. B. Fig. 2 — Classopoliis torosus (Reissinger) Couper 1958. Fig. 4 — Retimonocolpites sp. A. Fig. 5 — Manosulcites sp. A. Fig. 6 — Spore type A. Fig. 7 — Spore type B.

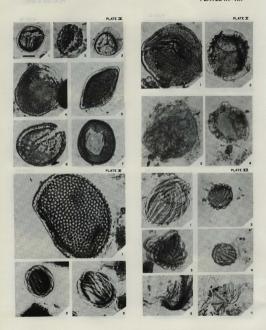
PLATES I-IV



PLATES V-VIII



PLATES IX - XII



PLATES XIII - XVI

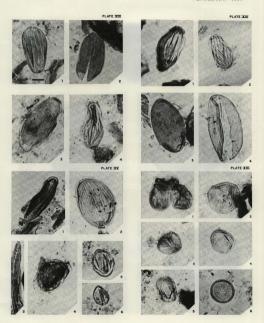


PLATE XVII

