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GEOLOGICAL SETTING

The Bom Sucesso region is located on the border of the Sao Francisco Craton, approximately 200 km south of Bel bortroate. The geology includes medium - and high-quade metamorphic and igneous rocks, with gneisses, amphibolites, migmatites, charrockites and granites. These have a probable Archema age and were reworked in the Transmazzonico cycle, with folding of the Bom Sucesso range, followed by granitic intrusions. The Bom Sucesso range corresponds to a shear root termoding NGO 1435 (MOMES & MULAMAM, 1937), QUORDERE, 1937). Morber to a shear root termoding NGO 1435 (MOMES & MULAMAM, 1937), QUORDERE, 1937). Morber zone. In this latter zone, Transmazonico rocks were recently suited at 1932-21 Ma by P. Vidal for France.

ORIENTATION AND DISTRIBUTION OF DIKES

The largest number of dikes occurs in the NM Archean block of the Bow Sucesso range. They are relatively scarce within Transamazônico granites, abundant in Archean granites and absent in the late Proteroodic São Jaão Del Rel Group. These dikes occur in very many directions. However, three main directions can be pointed out: 1) NIO[®]-AND[®] group: these are parallel to the Bow Sucesso range, appearing in large number on the NM side of the range; 2) NN-SE group; 3) ESS-MWN-NN INTO[®]-AND[®]-AND direction. Besides these, there are also a number of N-S and E-N diskes, associated with the E-W rapid trove.

The dikes have variable dimensions and thicknesses from a few decimeters up to 80 m, with an average of 10-30 m. The longest dikes (>10 m thick) are on the order of a few kilometers, reaching 30 km, in one case, in the Tabubes granite.

PETROGRAPHY

Doleritic textures dominate, with a mineralogy characterized by plagicclase, pyroxene and/or amphibole, quartz and rare biotite. Mafic minerals range from 25 to 85%. Narrow dikes and borders of wide dikes are fine grained. Some dikes are porphyritic.

The N110°-N120° group of dikes exhibits relict Ti-rich pyroxene, with irregular alteration to fine hornblende crystals. K-feldspar frequently forms myrmekitic and micrographic intercrysths with quarts.

In other dikes hornblende and actinolite appear in mafic aggregates of varied grain size. A 10-km-long porphyritic dike shows prisms 2-10 cm in diameter. These have pericline

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twinning and contain many acicular amphibole inclusions oriented according to plagicclase crystallographic directions. Plaglicclase in most dikes is labradorite and may be zoned.

Two chemical analyses indicate a tholelitic composition.

ACE

Tress dikes were formerly attributed to the Cretacous by ERICHEM (1929) and on the 1:1.000,000 scale map of Mhans Cerais (TEDEIRA & OSCIA & RAMMA, 1976). A/A dation TEDEIRA (1985) indicated a Precambrian age for dikes north of the Bom Sucesso region, despite a dispersion of dates.

Detailed field geology showed that the dike which cuts the Tabubes granite in the north is out by aplitic and pegmatitic veins related to late-stage crystallization of the granite. Therefore, they are probable synchronous with an age of 1930 Ms. These characteristics are also noted on the northern border of the Bom Sucesso range and in the E-M dike cutting the Tabubes granite.

In relation to the late Proterozoic folding of the São João Del Rei Group, it is clear that most dikes cross-cutting the Ritápolis granite were partially or completely schistified by this event. Such dikes are probably younger than the São João Del Rei Group.

The different number of dikes occurring in Precambrian rooks and in Transamazônico granites suggests the existence of pre-fransamazônico dikes. Although age dating of the dikes is not completely resolved, it is possible, however, that most of them are Transamazônico in age.

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